

SHIPBORNE SCATTEROMETER MEASUREMENTS OF GREAT LAKES ICE *

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ABSTRACT

The Jet Propulsion Laboratory C-band polarimetric scatterometer was used to measure radar backscatter of Great Lakes ice during the 1997 Great Lakes winter experiments. The scatterometer data are at the same frequency band, incident angles, and polarizations of operating satellite SARs such as RADARSAT and ERS or future multipolarization SARs such as ENVISAT and RADARSAT-II. During the experiments, in situ data for different ice types were obtained and accurate radar calibration measurements were conducted. We present backscatter results over the range of incident angles up to 60° for various ice types with different physical conditions, feature scale, thickness, snow cover, and concentration. The scatterometer data set is useful for the development of the Great Lakes ice mapping algorithm.

1.0 INTRODUCTION

Winter ice cover over the Great Lakes has a major impact on the regional climate, local commerce, and public safety (Assel et al., 1983). Millions of dollars in coal and ore shipping can be saved by extending the winter navigation season (Cooper et al., 1975). Ice is potentially harmful to the hydropower industry installations on the Niagara River (Assel et al., 1983). The number of days that ice cover exceeded 40% is an important input parameter to a whitefish recruitment forecast model in Lake Michigan (Brown et al., 1993). Ice cover duration is predicted by several General Circulation Models to reduce significantly under the CO₂ doubling scenario. Winter ecology may be strongly affected by the ice cover reduction (Vanderploeg et al., 1992). Ice jams not only impede navigation but also cause dangerous flooding (Daly, 1992). These scientific and practical applications require the monitoring of Great Lakes from satellite SAR with a good coverage and a high resolution (Leshkevich et al., 1995). In view of current and future satellite SARs, all operated at C band, we carried out experiments over Lakes Superior and northern Lake Michigan to measure C-band backscatter signatures of various ice types to develop ice mapping algorithms for the satellite SAR data.

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2.0 THE 1997 GREAT LAKES WINTER EXPERIMENTS

During the 1997 winter season, we conducted two experiments across the Straits of Mackinac and Lake Superior. The experimental campaign was coordinated into two expeditions on two different USCG Ice Breaker vessels, the Biscayne Bay in February and the Mackinaw (an Arctic-class ice breaker) in March. In these experiments, the Jet Propulsion Laboratory (JPL) polarimetric scatterometer was mounted on board the Biscayne Bay and the Mackinaw to measure the backscatter. The scatterometer was operated at C-band with the full polarization capability and incident angles from 0° to 60° . Thus, the results are applicable not only to RADARSAT SAR with the horizontal polarization, but also with ERS SAR with the vertical polarization and the future ENVISAT SAR with dual polarizations.

A Global Positioning System (GPS) receiver unit was used to record the locations along the ship routes when the scatterometer data were taken. The GPS data were downloaded to the computer that controlled the radar in near-real time. The data were plotted out on a map of the Great Lakes to show the locations of the ice types where radar data were obtained. The computer internal clock was synchronized with the GPS time and both radar data and GPS data were time tagged so that they can be correlated. An anemometer was used near the radar to measure in-situ wind. The ship instruments also provided location, wind data, and temperature data.

During the experiments, accurate calibration measurements were conducted to calibrate the scatterometer data. A trihedral corner reflector of known radar cross section was used for this purpose. The corner reflector was aligned with the radar pointing direction with a laser pointer. Another calibration target, a metallic sphere, was used to obtain additional data to validate the radar calibration. During the experiment on board the Biscayne Bay, the calibration measurements were taken while the ship docked near De Tour Village in Michigan. In the Mackinaw experiment, the calibration targets were setup on an open area at the back of the ship and microwave absorbers were used to avoid unwanted signals from reflections and multipath effects.

Ice truth data for different ice types were obtained. Ice thickness and snow thickness were measured. Photographs showing the layering structure of snow ice and lake ice with different amount of air inclusions were taken. Wind speed and temperature were recorded. We also took a number of snow thickness measurements over an area at a location during the Biscayne Bay experiment to determine the range of snow thickness distribution. Photographs were taken to estimate the surface roughness condition and snow coverage. For each set of radar measurements, photographs were taken both in the near range coincident to the area of the radar footprint and in the far range for an overall picture of the ice type.

3.0 BACKSCATTER SIGNATURE RESULTS

In the 1997 experimental campaign, we acquired about 2000 radar data files for many ice types under different physical and environmental conditions. For each ice type, we

obtained radar data from 0° to 60° incident angles. For each incident angles, the measurements covered several different samples in azimuth directions scanned over the targeted area. For each antenna look direction, data were coherently averaged by 20 times over 401 samples in the frequency domain with 1 GHz bandwidth. The data included both magnitudes and phases of the scattering matrix for all combinations in the linear polarization basis. The data are processed and calibrated into polarimetric backscattering coefficients.

We acquired backscatter signatures of various ice types with different physical condition, feature scale, thickness, snow cover, and concentration. Some of the ice types are pancake ice with complicated structures, finely crushed ice refrozen into a matrix of lake ice with a very thin snow cover, and plate ice with angular polygon shape with a larger scale compared to the rounded pancake ice. A typical ice type over the Great Lakes is snow ice (white ice) over lake ice (black ice) with various thickness. We encountered a number of rubble ice fields produced by strong wind and wave actions which broke and pushed the ice together. Broken layers of ice, whose total thickness reached several meters, were seen in this highly deformed ice type. New black ice of a few inch thick was transparent (almost no bubble) with very smooth upper and lower surfaces. Snow cover from a dusting condition to a thick layer was observed.

Different ice types can have different backscattering behaviors. For example, pancake ice, consisting of broken pieces of ice with rounded shape and a raised rim around the edge, has backscatter σ_{vv} decreasing quickly with incident angle. However, unlike the typical rough surface scattering behavior, the copolarized ratio $\gamma = \sigma_{vv}/\sigma_{hh}$ is negative or close to zero. The crosspolarized ratio $e = \sigma_{hv}/\sigma_{hh}$ of this ice type is small at large incident angles, and the correlation coefficient ρ between the horizontal and vertical is small at large incident angles.

Measured backscatter data of typical snow covered lake ice indicate that the horizontal backscatter is larger than the vertical backscatter especially at larger incident angles. This backscatter property of lake ice is different from that of most sea ice types whose vertical return is usually larger than the horizontal one. For future ENVISAT SAR data with both polarizations, this result is useful to distinguish the fresh water ice from open water, which has $\sigma_{vv} > \sigma_{hh}$. Radar data of the typical lake ice type taken in March along the ship track show that C-band waves can propagate more than 1 m in the ice. For deformed ice in a rubble field, the backscatter is very strong across the range of incident angles. Black ice with a thin snow cover has low backscatter with a strong decreasing gradient in incident angle. The polarimetric scatterometer data set is useful for the development of the Great Lakes ice mapping algorithm. Furthermore, this data set can be used to determine which ice type can be observed for a given set of operating parameters of a satellite SAR, such as the system noise floor, polarizations, and incident angles.

6.0 REFERENCES

- R. A. Assel, F. H. Quinn, G. A. Leshkevich, and S. J. Bolsenga, *Great Lakes Ice Atlas*, NOAA, Great Lakes Environmental Research Laboratories, Ann Arbor, Mich., 1983.

- R. W. Brown, W. W. Taylor, and R. A. Assel, "Factors affecting the recruitment of lake whitefish in two areas of northern Lake-Michigan," *J. Great Lakes Res.*, vol. 19, no. 2, pp. 418-428, 1993.
- D. W. Cooper, R. A. Mueller, and R. J. Schertler, "Remote profiling of lake ice using an S-band short-pulser radar aboard an all terrain vehicle," *NASA Tech. Mem.*, NASA TM X-71808, 1975.
- S. F. Daly, "Observed ice passage from Lake Huron into the St-Clair river," *J. Great Lakes Res.*, vol. 18, no. 1, pp. 61-69, 1992.
- G. A. Leshkevich, W. Pichel, P. Clemente-Colon, R. Carey, and G. Hufford, "Analysis of coastal ice cover using ERS-1 SAR data," *Int. J. Rem. Sens.*, vol. 16, no. 17, pp. 3459-3479, 1995.
- H. A. Vanderploeg, S. J. Bolsenga, G. L. Fahnenstiel, J. R. Liebig, and W. S. Gardner, "Plankton ecology in an ice-covered bay of Lake Michigan: utilization of a winter phytoplankton bloom by reproducing copepods," *Hydrobiologica*, 243-244, 175-183, 1992.

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